

Coral Energy Reserves and Calcification in a High-CO₂ World at Two Temperatures

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Abstract

Rising atmospheric CO₂ concentrations threaten coral reefs globally by causing ocean acidification (OA) and warming. Yet, the combined effects of elevated pCO₂ and temperature on coral physiology and resilience remain poorly understood. While coral calcification and energy reserves are important health indicators, no studies to date have measured energy reserve pools (i.e., lipid, protein, and carbohydrate) together with calcification under OA conditions under different temperature scenarios. Four coral species, *Acropora millepora*, *Montipora monasteriata*, *Pocillopora damicornis*, *Turbinaria reniformis*, were reared under a total of six conditions for 3.5 weeks, representing three pCO₂ levels (382, 607, 741 μatm), and two temperature regimes (26.5, 29.0°C) within each pCO₂ level. After one month under experimental conditions, only *A. millepora* decreased calcification (−53%) in response to seawater pCO₂ expected by the end of this century, whereas the other three species maintained calcification rates even when both pCO₂ and temperature were elevated. Coral energy reserves showed mixed responses to elevated pCO₂ and temperature, and were either unaffected or displayed nonlinear responses with both the lowest and highest concentrations often observed at the mid-pCO₂ level of 607 μatm. Biweekly feeding may have helped corals maintain calcification rates and energy reserves under these conditions. Temperature often modulated the response of many aspects of coral physiology to OA, and both mitigated and worsened pCO₂ effects. This demonstrates for the first time that coral energy reserves are generally not metabolized to sustain calcification under OA, which has important implications for coral health and bleaching resilience in a high-CO₂ world. Overall, these findings suggest that some corals could be more resistant to simultaneously warming and acidifying oceans than previously expected.

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Introduction

Anthropogenic climate change threatens many marine ecosystems today, and coral reefs are among the most sensitive to current changes in ocean biogeochemistry [1,2]. Rising atmospheric carbon dioxide (CO₂) concentrations have already caused an increase of 0.6°C in the average temperature of the upper layers of the ocean over the past 100 years [3], and about one third of all anthropogenic CO₂ has been absorbed by the ocean, causing ocean acidification (OA) [4,5]. Since scleractinian corals are calcifying organisms that already live close to their upper thermal tolerance limits [6], both ocean warming and acidification severely threaten their survival and role as reef ecosystem engineers [1,7].

The uptake of anthropogenic CO₂ by the ocean changes the carbonate chemistry of seawater by increasing proton (H⁺) and bicarbonate (HCO₃[−]) concentrations, while at the same time decreasing the concentration of carbonate (CO₃^{2−}). Consequently, seawater pH (i.e., $-\log[H^+]$) and the saturation state with respect to aragonite ($\Omega_{\text{arag}} = [Ca^{2+}][CO_3^{2-}]/K_{\text{sp}}$ with K_{sp} being the ionic product of [Ca²⁺] and [CO₃^{2−}] under solution-mineral equilibrium) decreases. As aragonite is the form of calcium carbonate (CaCO₃) precipitated by modern corals, this process compromises marine calcification [8–10]. Over the past century, Ω_{arag} in the tropics has decreased from 4.6 to 4.0 [8] and is expected to decrease to 2.5–3.0 by the year 2100 [1,8,11]. Further, it has been estimated that scleractinian calcification rates may drop by up to 35%–40% by the end of this century [8,12].

Coral calcification typically decreases in response to experimentally reduced seawater pH [13–24] but not always [14,15,25–31]. Seawater temperature also influences calcification [32–36], resulting in potentially interactive effects of temperature and OA on coral calcification. For example, negative effects of elevated seawater $p\text{CO}_2$ on calcification are often exacerbated when temperature is simultaneously increased [20,27,30], suggesting a synergistic interactive effect. However, this is not always observed [19,28,31] and in one study even the opposite was shown [37]. Clearly, further studies are required to gain a better understanding of the interactive effects of elevated temperature and $p\text{CO}_2$ on coral calcification and its resistance to OA.

Much less is known about how combined OA and warming will influence other aspects of coral physiology such as energy reserves and tissue biomass. If calcification becomes energetically more costly under elevated $p\text{CO}_2$ due to a decreased aragonite saturation state [38–40], then the extra energy needed to maintain calcification might be drawn from one or more of the following sources: 1) Coral energy reserves (i.e., lipids, protein, carbohydrates), 2) Enhanced endosymbiotic algal production due to CO_2 fertilization [41], and 3) Increased heterotrophy (i.e., zooplankton, particulate and/or dissolved organic carbon) [28,42]. These responses may be even more extreme with the simultaneous increases in seawater temperature because tissue biomass, energy reserves, and endosymbiotic algal density are typically lowest when temperature (and irradiance) is highest on seasonal timescales [43,44,45] and under bleaching scenarios [46,47,48].

Although tissue biomass and energy reserves are important indicators of coral health [46,47] and play a significant role in promoting resilience to bleaching [49], no studies to date have measured all three energy reserve pools (i.e., lipid, protein, and carbohydrate) under OA conditions at elevated temperature. While protein concentrations were either unaffected [30,31] or increased in response to elevated $p\text{CO}_2$ alone [13,50], the effects of OA, or OA plus elevated temperature, on coral lipids and carbohydrates are unknown. Studies specifically addressing all three energy reserve pools are needed to get a better understanding of how OA affects coral energetics and their overall resistance to future climate change.

Finally, the algal endosymbiont (*Symbiodinium* sp.) provides healthy corals with up to 100% of their daily metabolic energy demand via photosynthesis [51]. If algal productivity is enhanced under OA due to CO_2 fertilization [41], this might help maintain calcification rates and/or energy reserves under OA as energetic costs for calcification increase. Further, *Symbiodinium* sp. exhibit high sensitivity to elevated seawater temperature [52]. Thus, it is important to monitor endosymbiont and chlorophyll *a* concentrations in studies manipulating both $p\text{CO}_2$ and temperature.

Here, we studied the single and interactive effects of $p\text{CO}_2$ (382, 607, 741 μatm) and temperature (26.5 and 29.0°C) on coral calcification, energy reserves (i.e., lipid, protein, and carbohydrate), chlorophyll *a*, and endosymbiont concentrations in 4 species of Pacific coral with different growth morphologies. It was hypothesized that 1) calcification and energy reserves decrease in response to elevated $p\text{CO}_2$ and elevated temperature, 2) decreases are larger when $p\text{CO}_2$ and temperature are elevated simultaneously, and 3) that physiological responses are species-specific. We show that only one of the four coral species studied here decreased calcification in response to average ocean acidification levels expected by the second half of this century (741 μatm), even when combined with elevated temperature (+2.5°C). Further, we show for the first time that energy reserves were largely not metabolized in order to sustain calcification under elevated $p\text{CO}_2$ and temperature, suggesting that some coral species will be more

resistant to combined ocean acidification and warming than previously expected.

Materials and Methods

Experiment

Six parent colonies of *Acropora millepora*, *Pocillopora damicornis*, *Montipora monasteriata*, and *Turbinaria reniformis* were purchased from Reef Systems Coral Farm (New Albany, Ohio, USA) which is a CITES permit holder. The parent colonies were specifically collected for this experiment from 3–10 m in northwest Fiji (17°29'19"S, 177°23'39"E) in April 2011. Colonies of the same species were collected at least 10 m apart to increase the probability that different genotypes of the same species were selected. All colonies were shipped to Reef Systems Coral Farm and maintained in recirculating indoor aquaria with natural light (greenhouse, 700–1000 $\mu\text{mol quanta m}^{-2} \text{s}^{-1}$) and commercially available artificial seawater (Instant Ocean Reef Crystals) for 2.5 months until the start of the experiment.

From April 22 - May 19, 2011, six fragments were collected from each parent colony and mounted on PVC tiles for a total of 144 fragments (4 species \times 6 colonies \times 6 fragments; Fig. 1). Starting on June 19, 2011, corals were gradually acclimated to a custom-made artificial seawater (ESV Aquarium Products Inc.), which was designed to mimic the chemical composition and alkalinity of natural reef seawater. On July 8 and 9, 2011, all 144 fragments were transferred to the experimental recirculating indoor aquaria with artificial light (Tek Light T5 actinic lights, 275 $\mu\text{mol quanta m}^{-2} \text{s}^{-1}$, 9:15 hrs light:dark cycle) and allowed to acclimate to the artificial light conditions for 10 days under ambient seawater conditions (i.e., 26.5°C and $p\text{CO}_2$ of 382 μatm). Photosynthesis to irradiance (P/E) curves performed on *Acropora millepora* showed that photosynthesis was fully saturated at these light levels. Due to logistical reasons, P/E curves were not performed on the other species.

For each of the 6 treatments, the recirculating tank system consisted of one 905 L sump and six aquaria of 57 L each. One fragment per parent colony per species was put in one of the 6 aquaria in each system such that there were a total of 4 fragments (one of each species) in each aquarium, and each parent colony of each species was represented in each system. By placing the same genotypes in each treatment, genotypic variation between treatments was minimized and our ability to detect treatment effects was optimized. Replication of treatments and independent tanks within treatments was not possible due to the complexity and cost of operating tanks under modified pH conditions. While this is, strictly speaking, a pseudo-replicated design [53], the disadvantages of this design are outweighed by the advantages of being able to simultaneously manipulate six combinations of temperature and pH. To optimize the experimental design conditions, coral fragments were rotated daily within tanks and every 3 days among tanks within each system to minimize any tank or positional effects within each system. Further, tanks were cleaned every three days, and great care was taken to ensure similar conditions across treatments except for carbonate chemistry and temperature.

Experimental treatments were assigned to each system as follows: 26.5°C and 382 μatm , 26.5°C and 607 μatm , 26.5°C and 741 μatm , 29.0°C and 382 μatm , 29.0°C and 607 μatm , and 29.0°C and 741 μatm (Fig 1). The three $p\text{CO}_2$ levels—382, 607, and 741 μatm —were designed to represent present day $p\text{CO}_2$, and two $p\text{CO}_2$ levels expected by the second half of the 21st century, respectively. The control temperature (26.5°C) represents the current average annual temperatures in Fiji (<http://www.ospo>).

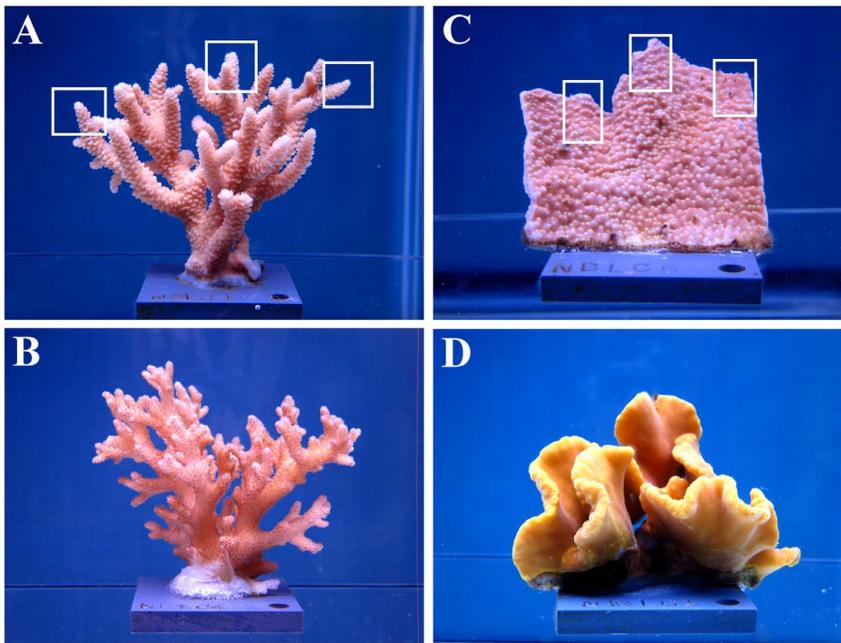


Figure 1. Photos of representative coral fragments from (a) *Acropora millepora*, (b) *Pocillopora damicornis*, (c) *Montipora monasteriata*, and (d) *Turbinaria reniformis*. Rectangles indicate subsamples taken from each fragment for lipid, protein/carbohydrate, and tissue biomass analyses. The remaining tissue was airbrushed for chlorophyll *a* and endosymbiont density measurements.
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noaa.gov/Products/ocean/index.html), whereas 29.0°C represents the upper limit of current summer temperatures but is still below the bleaching threshold at that location. Therefore, the 26.5°C and 382 μatm treatment served as control. The experiment lasted for 24 days from July 19-August 12, 2011.

Temperature was controlled by titanium aquarium heaters submerged in each system sump (Aqua Medic) and connected to a digital control system (Neptune Systems Apex AquaController). Temperature loggers (Onset Hobo Pro v2) were placed in each sump and recorded temperature every 5 minutes. Seawater μCO_2 was controlled by bubbling in pure CO_2 , CO_2 -free air, or ambient air delivered by an outdoor air pump (Sweetwater, Aquatic Eco-Systems Inc.) into each system sump. CO_2 -free air was achieved by moving ambient air through CO_2 -scrubbers consisting of a 1.5 m long tube (10 cm diameter) filled with soda lime (SodaSorb HP). Supply of all gases was controlled via a pH stat system using custom designed software (KSgrowstat, written by K. Oxborough, University of Essex). Seawater pH was measured every 5 seconds by microelectrodes (Thermo Scientific Orion Ross Ultra pH glass electrode), which were calibrated daily.

For the elevated temperature (29.0°C) treatments, temperature was gradually increased over several days until the desired temperature was reached. For the medium (607 μatm) and high (741 μatm) μCO_2 treatments, μCO_2 was gradually increased over several days starting from 382 μatm until the final μCO_2 was achieved. Recirculating seawater flow rate was 210–230 l/hour and little pumps (Accela Powerheads) created additional water circulation within each aquarium. A quarter of the entire water volume of each treatment system was exchanged every 3 days. Non-carbonate ceramic filter media (MarinePure High Performance Biofilter Media, CerMedia) were placed in the sumps to filter the water. Tanks were cleaned every 3 days or as needed.

Since healthy corals *in situ* can acquire up to 46% of their daily metabolic energy demands by feeding on zooplankton [48,54], corals were fed every three days with 48 h old brine shrimp nauplii

(*Artemia* sp., Carolina Biological Supply). Corals were allowed to acclimate to the dark for 30 min before feeding was conducted. They were fed for one hour in separate, partially submerged plastic containers containing water from their respective treatment, and at a concentration of approximately 1 brine shrimp nauplii ml^{-1} which is representative of zooplankton concentrations on natural Pacific reefs [55]. At the end of the hour, brine shrimp nauplii remained in the feeding chambers indicating that the corals had not captured all brine shrimp nauplii available to them. Following feeding, the corals were placed back in their respective aquaria and the feeding container water discarded so as not to introduce brine shrimp into the recirculating systems.

Monitoring of seawater chemistry during the experiment

Temperature and salinity were measured daily (YSI 63), and salinity was adjusted daily to 35 ppt. Daily water samples were taken using screw-top high-density polyethylene bottles for pH and alkalinity analyses. After equilibration at 25°C in a recirculating water bath (30 min), sample pH_{NBS} was measured with an Orion® Ross glass electrode (precision 0.01 pH units) [56], which was calibrated daily at 25°C. Total alkalinity (TA) was titrated with HCl on the same samples using an AS-ALK2 (Apollo SciTech Inc.) alkalinity titrator [57] (precision 0.1%). The HCl solution was calibrated with Certified Reference Material (CRM) from A.G. Dickson (Scripps).

Treatment $x\text{CO}_2$ (dry air), aragonite saturation state (Ω_{arag}), and pH_{T} were calculated using the program CO2SYS [58] based on measured pH_{NBS} and alkalinity at the respective temperature. $x\text{CO}_2$ was converted to μCO_2 using the equation in Weiss *et al.* [59]. Carbonate dissociation constants were taken from Millero *et al.* [60]. In addition, a custom-made CO_2 analyzer based on a LICOR 820 was used weekly to crosscheck with calculated sump $x\text{CO}_2$ values according to methods by Wang & Cai [56], and indicated good agreement of measured and calculated values ($r^2 = 0.97$, $n = 66$).

Laboratory analyses

Calcification. Net calcification was determined using the buoyant weight technique [61]. Each coral fragment was buoyantly weighed at the beginning, middle (after 11 experimental days), and at the end of the experiment (after 23 experimental days). As such, it was possible to assess if calcification rates varied during the experiment. Daily calcification rates were calculated as the difference between initial, middle, and final weights, divided by the respective number of days elapsed, and standardized to surface area (see below).

For tissue analyses, corals were frozen at -80°C and a total of three branch tips or growing edge pieces were saved from each fragment for lipid, protein/carbohydrate, and tissue biomass analyses, respectively (Fig. 1). The remaining tissue was airbrushed for chlorophyll *a* and endosymbiont density measurements.

Chlorophyll *a* and endosymbiont density. Coral tissue was stripped off the coral skeleton with a waterpik [62] containing 40 ml of synthetic seawater (Instant Ocean). The endosymbionts were isolated from the host tissue via centrifugation and then resuspended in 10 ml of synthetic seawater. For chlorophyll *a* concentrations, 1 ml of this algal suspension was pelleted and the cells lysed in 1 ml of 4°C methanol using a bead-beater for 60 seconds. Samples were then immediately placed on ice and allowed to extract for one hour in the dark. Samples were centrifuged to remove cellular debris and measured spectrophotometrically ($\lambda = 652, 665 \text{ \& } 750$) on a 96-well plate reader. The equations for chlorophyll *a* in methanol described by Porra *et al.* [63], along with path length correction [64], were used to calculate chlorophyll *a* concentrations (pg/cell), and were then standardized to surface area (see below). Another 1 ml subsample of the algal suspension was preserved with 10 μl of 1% glutaraldehyde solution for endosymbiont quantification, which was calculated using 6 independent replicate counts on a hemocytometer, using a Nikon microphot-FXA epifluorescent microscope at $100\times$ magnification. Photographs were analyzed through Image J using the analyze particles function.

Energy reserves and tissue biomass. For all energy reserve and tissue biomass measurements, only branch tips or samples with a growing edge were used. While tissue composition may vary across the surface of a coral [65], this approach was used to allow for comparison with previously published studies [46,47,66]. Soluble lipids (referred to hereafter simply as lipids) were extracted from a whole, ground coral sample (skeleton + animal tissue + algal endosymbiont) in a 2:1 chloroform:methanol solution for 1 hour [46,66] washed in 0.88% KCl followed by 100% chloroform and another wash with 0.88% KCl. The extract was dried to constant weight under a stream of pure nitrogen (UPH grade 5.0) and standardized to the ash-free dry weight.

Animal soluble protein and carbohydrate (referred to hereafter simply as protein and carbohydrate, respectively) were extracted from grinding a whole second branch tip of the same fragment [46]. Briefly, Milli-Q water was added to the ground coral sample and the resulting slurry was sonicated (5 min) and then centrifuged twice (5000 rpm, 10 min) to separate the animal tissue from the skeleton and endosymbiotic algae. Protein and carbohydrate was extracted from the animal tissue only. One subsample of this animal tissue slurry was used for protein extraction using the bicinchoninic acid method [67] with bovine serum albumin as a standard (Pierce BCA Protein Assay Kit). A second subsample was used for carbohydrate quantification using the phenol-sulfuric acid method [68] with glucose as a standard. Soluble animal protein and carbohydrate concentrations were standardized to the ash-free dry weight.

Tissue biomass was measured by drying a third branch tip of whole coral sample (skeleton + animal tissue + algal endosymbiont) to constant dry weight (24 hours, 60°C) and burning it (6 hours, 450°C). The difference between dry and burned weight was the ash free dry weight which was standardized to the surface area of this branch tip.

Surface area. Surface area of plating *M. monasteriata* and *T. reniformis* fragments was determined using the aluminum foil technique [69], whereas surface area of branching *A. millepora* and *P. damicomis* fragments was determined using the single wax dipping technique [70,71] after the tissue had been removed. Natural wooden blocks of varying sizes and shapes were used as calibration standards [71]. Wax dipping was conducted using household paraffin wax (Gulf Wax, Royal Oak Enterprises) heated to 65°C . Dried coral skeletons and wooden calibration standards were maintained at room temperature prior to weighing.

Statistical analyses

Three-way mixed-model analyses of variance (ANOVA) tested the effects of $p\text{CO}_2$, temperature, and parent colony on calcification rates in the first and second half of the experiment, chlorophyll *a*, algal endosymbiont density, lipid, protein, carbohydrate, and tissue biomass. Temperature and $p\text{CO}_2$ were fixed and fully crossed, whereas parent colony was a random factor. The ANOVAs were run for each species separately. All data were normally distributed according to plots of residuals versus predicted values for each variable, or transformed to meet the condition of normality. Outlier values greater than 3 times the interquartile range were excluded. Post hoc Tukey tests were performed when main effects were significant ($p \leq 0.05$). A posteriori slice tests (e.g., tests of simple effects) [72] determined if the ambient (26.5°C) and elevated (29.0°C) temperature treatment averages significantly differed within each $p\text{CO}_2$ level. Bonferroni corrections were not applied [73,74], therefore significant model p -values > 0.0016 (0.05/32 tests) should be interpreted with caution. Statistical analyses were performed using SAS software, Version 9.2 of the SAS System for Windows.

Results

All corals appeared healthy throughout the experiment. No visible paling and no mortality occurred. The average seawater temperature, pH_T , $p\text{CO}_2$, saturation state, and total alkalinity for all six treatments are summarized in Table 1.

Calcification

In *Acropora millepora*, calcification rates during the first half of the experiment were overall unaffected by both temperature ($p = 0.36$) and $p\text{CO}_2$ ($p = 0.79$) (Fig. 2a; Table S1). However, at the highest $p\text{CO}_2$ level (741 μatm) calcification was 43% lower at 29.0°C than at 26.5°C (Fig. 2a). During the second half of the experiment, calcification rates were significantly affected by $p\text{CO}_2$ ($p = 0.001$) but not temperature ($p = 0.42$), and were lower by 53% at the highest compared to the lowest $p\text{CO}_2$ level (Fig. 2b; Table S1).

In *Pocillopora damicomis*, a significant interaction of $p\text{CO}_2$ and temperature ($p < 0.001$) was observed for calcification rates during the first half of the experiment (Fig. 2c, Table S1). During the second half of the experiment, calcification rates were generally unaffected by temperature ($p = 0.06$) and $p\text{CO}_2$ ($p = 0.07$). However, at ambient seawater $p\text{CO}_2$ (382 μatm) corals kept at elevated temperature (29.0°C) calcified 91% more compared to those kept at 26.5°C (Fig. 2d, Table S1).

Calcification rates of *Montipora monasteriata* were affected by temperature ($p = 0.04$) but not $p\text{CO}_2$ ($p = 0.42$) during the first half

Table 1. Average conditions for each of the 6 treatments representing three pCO₂ levels at two temperature regimes (ambient, elevated = ambient + 2.5°C).

	400 ppm target		600 ppm target		800 ppm target	
	ambient temp.	elevated temp.	ambient temp.	elevated temp.	ambient temp.	elevated temp.
Temp.(°C)	26.45±0.01	29.31±0.02	26.37±0.01	28.53±0.02	26.61±0.01	28.93±0.02
pH _T	8.07±0.01	8.04±0.01	7.90±0.01	7.89±0.01	7.83±0.01	7.81±0.01
pCO ₂ (iatm)	364.31±9.69	400.62±16.83	598.37±18.50	616.08±24.24	732.04±22.37	749.63±26.21
TA (μmol kg ⁻¹)	2269.4±10.84	2270.1±11.15	2303.8±9.34	2288.3±10.43	2306.3±10.64	2304.5±9.08
Ω _{arag}	3.69±0.07	3.79±0.09	2.75±0.05	2.91±0.05	2.40±0.06	2.52±0.06

Mean ± 1 SE are shown. Sample size was 25 for all measurements. Temp. = Temperature.
doi:10.1371/journal.pone.0075049.t001

of the experiment (Fig. 2c–f, Table S1), with corals calcifying 18% more at elevated compared to ambient temperature. This was largely driven by significant temperature differences at both 382 and 607 μatm but not 741 μatm. During the second half of the experiment, calcification rates were unaffected by both temperature ($p = 0.82$) and pCO₂ ($p = 0.14$).

In contrast, calcification rates of *Turbinaria reniformis* during both first and second half of the experiment (Fig. 2g–h, Table S1) did not respond to changes in seawater temperature ($p = 0.45$ and 0.17) or pCO₂ ($p = 0.36$ and 0.09). Notably, the two plating species (*M. monasteriata* and *T. reniformis*) calcified more than twice as fast as the two branching species (*A. millepora* and *P. damicornis*).

Chlorophyll a and endosymbiont density

The chlorophyll *a* concentrations of *A. millepora* were significantly affected by pCO₂ ($p < 0.001$) but not temperature ($p = 0.054$), with concentrations being 51% lower at 607 μatm than at either 382 or 741 μatm (Fig. 3a, Table S2). Endosymbiont densities were not affected by either seawater temperature

($p = 0.07$) or pCO₂ ($p = 0.03$ but overall model $p = 0.24$) (Fig. 3b, Table S2).

In *P. damicornis*, a significant interaction of temperature and pCO₂ was observed for both chlorophyll *a* concentrations and endosymbiont densities ($p < 0.001$ and $p = 0.02$, respectively) (Fig. 3c–d, Table S2). When temperature was elevated, chlorophyll *a* concentrations were higher by 19% and 67%, respectively, at both 382 and 741 μatm. At 607 μatm, endosymbiont densities decreased by 36% at 29°C compared to concentrations at ambient temperature.

In *M. monasteriata*, a significant interaction of seawater temperature and pCO₂ was observed for chlorophyll *a* concentrations ($p = 0.01$) (Fig. 3e, Table S2), with concentrations being 45% and 30% lower at elevated compared to ambient temperature, respectively, under both 382 and 741 μatm conditions. Endosymbiont densities were significantly affected by both temperature ($p < 0.001$) and pCO₂ ($p = 0.01$) but the interaction term was not significant ($p = 0.38$) (Fig. 3f, Table S2). Densities were 32% lower at elevated compared to ambient temperature, and were lowest

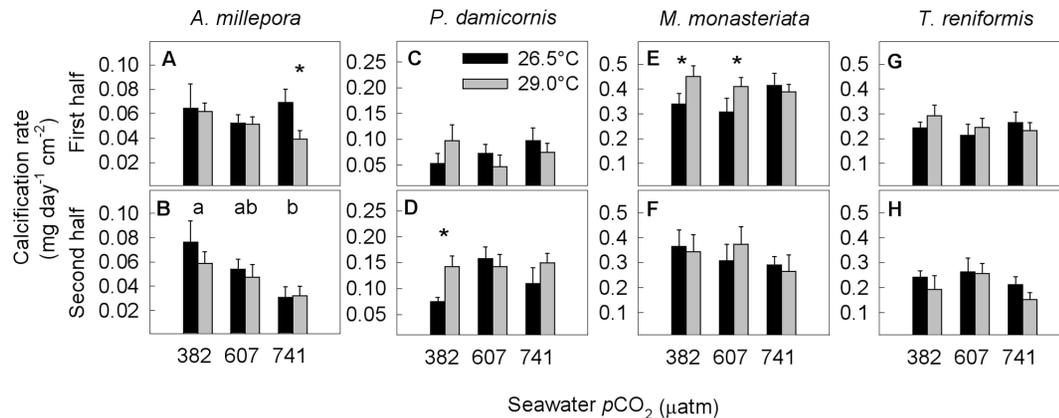


Figure 2. Average daily calcification rate during the first and the second half of the experiment for (a, b) *Acropora millepora*, (c, d) *Pocillopora damicornis*, (e, f) *Montipora monasteriata*, and (g, h) *Turbinaria reniformis*. Averages ± 1 SE are shown for three pCO₂ levels and two temperature regimes (26.5, 29.0°C). Asterisks indicate significant differences between 26.5 and 29.0°C within a given pCO₂ level (determined by a posteriori slice tests). The letters a and b indicate results of the post hoc Tukey tests when there was a significant pCO₂ effect. Sample sizes ranged between 5 and 6. Statistical details can be found in Table S1.
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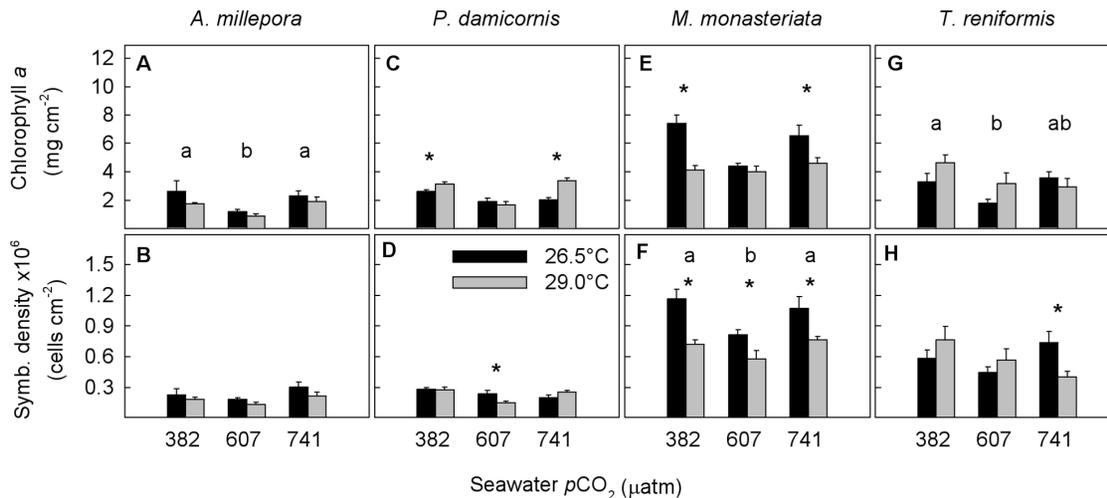


Figure 3. Average chlorophyll a concentrations and symbiont density for (a, b) *Acropora millepora*, (c, d) *Pocillopora damicornis*, (e, f) *Montipora monasteriata*, and (g, h) *Turbinaria reniformis*. Averages \pm 1 SE are shown for three pCO₂ levels and two temperature regimes (26.5, 29.0°C). Asterisks indicate significant differences between 26.5 and 29.0°C within a specific pCO₂ level (determined by a posteriori slice tests). The letters a and b indicate results of the post hoc Tukey tests when there was a significant pCO₂ effect. Sample sizes ranged between 5 and 6. Statistical details can be found in Table S2. doi:10.1371/journal.pone.0075049.g003

overall at 607 μatm (−25%) compared to the other two pCO₂ levels.

Chlorophyll a concentrations of *T. reniformis* were affected by pCO₂ ($p=0.03$) but not temperature ($p=0.11$), with concentrations being 38% lower at 607 compared to 382 μatm (Fig. 3g, Table S2). Endosymbiont densities were not affected by either temperature ($p=0.90$) or pCO₂ ($p=0.21$) but were lower by 45% at elevated compared to ambient temperature under 741 μatm conditions (Fig. 3h, Table S2).

Energy reserves and tissue biomass

Lipid concentrations of *A. millepora* were affected by seawater pCO₂ ($p=0.01$) but not temperature ($p=0.053$), with concentrations being 28% and 21% higher at 607 and 741 μatm, respectively, compared to concentrations at ambient pCO₂ (Fig. 4a, Table S3). A significant interaction of seawater pCO₂ and temperature was observed for protein concentrations ($p=0.01$) (Fig. 4b, Table S3). Carbohydrate concentrations were affected by both temperature ($p=0.02$) and pCO₂ ($p=0.01$) but the interaction term was not significant ($p=0.85$) (Fig. 4c, Table S3). Across all pCO₂ treatments, carbohydrate concentrations were 18% lower at 29.0°C compared to 26.5°C, and 41% lower at 607 μatm than at 741 μatm. Tissue biomass was unaffected by changes in seawater temperature ($p=0.99$) and pCO₂ ($p=0.07$) (Fig. 4d, Table S3).

In *P. damicornis*, lipid concentrations were affected by seawater pCO₂ ($p=0.01$) but not temperature ($p=0.53$) (Fig. 4e, Table S3), with concentrations being 41% and 18% higher at 607 and 741 μatm, respectively, compared to concentrations at ambient pCO₂ (Fig. 4e, Table S3). Neither protein, nor carbohydrate concentrations or tissue biomass were affected by seawater temperature ($p=0.63, 0.88, 0.33$, respectively) and pCO₂ ($p=0.52, 0.35, 0.41$, respectively) (Fig. 4f–h, Table S3).

The lipid concentrations of *M. monasteriata* were unaffected by both seawater temperature ($p=0.38$) and pCO₂ ($p=0.23$) (Fig. 4i, Table S3). A significant interaction of temperature and pCO₂ was observed for protein concentrations ($p<0.001$): at 382 μatm, they decreased (−27%) at elevated compared to ambient temperature,

whereas at 607 and 741 μatm, they increased (+36% and +60%, respectively) (Fig. 4j, Table S3). Carbohydrate concentrations were affected by seawater temperature ($p=0.02$) but not pCO₂ ($p=0.36$) (Fig. 4k, Table S3), and the concentrations were 25% higher at elevated than at ambient temperature under 741 μatm conditions. Tissue biomass was also unaffected by both seawater temperature ($p=0.78$) and pCO₂ ($p=0.12$) (Fig. 4l, Table S3).

In *T. reniformis*, none of the measured energy reserve pools responded to changes in seawater temperature and pCO₂ (Fig. 4m–o, Table S3). Tissue biomass was also unaffected by both temperature ($p=0.62$) and pCO₂ ($p=0.58$), but was 21% lower at 29.0°C compared to 26.5°C at 382 μatm (Fig. 4p, Table S3).

Effects of parent colony

Parent colony was a significant effect in many of the measured variables, but no single parent colony or group of specific parent colonies was consistently different from all other parent colonies in any of the species studied (Tables S1–S3).

Discussion

Coral calcification has been predicted to decrease dramatically by the end of this century, thus threatening the existence of coral reefs in the future. Although the response of coral calcification is not uniform across species, most studies have found that calcification decreases with increasing seawater pCO₂ [75–77]. Here, we show that only one of the four Pacific coral species studied here decreased calcification in response to average ocean acidification levels expected by the second half of this century (741 μatm), even when combined with elevated temperature (+2.5°C). Further, we investigated for the first time the effects of OA on coral energy reserves and show that they were largely not metabolized in order to sustain calcification under elevated pCO₂ and temperature.

Acropora millepora was the only coral out of the four species studied here that decreased calcification rates in response to OA (Fig. 2b). While calcification rates were not affected by elevated pCO₂ and/or temperature during the first half of the experiment, they declined by 53% during the second half of the experiment

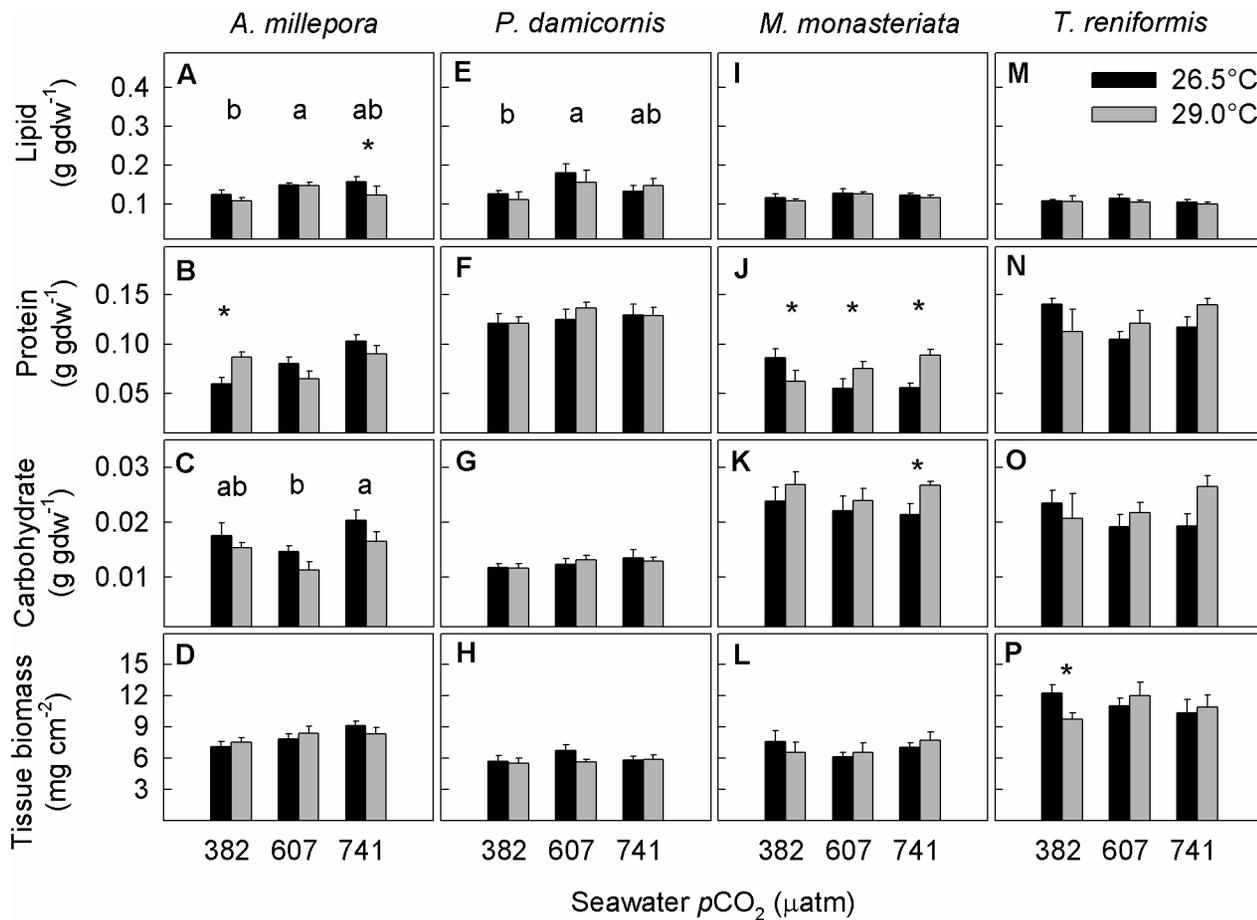


Figure 4. Average lipid, protein, carbohydrate concentrations, and tissue biomass of (a–d) *Acropora millepora*, (e–h) *Pocillopora damicornis*, (i–l) *Montipora monasteriata*, and (m–p) *Turbinaria reniformis*. Averages \pm 1 SE are shown for three pCO₂ levels and two temperature regimes (26.5, 29.0°C). Asterisks indicate significant differences between 26.5 and 29.0°C within a specific pCO₂ level (determined by a posteriori slice tests). The letters a and b indicate results of the post hoc Tukey tests when there was a significant pCO₂ effect. Sample sizes ranged between 4 and 6. Statistical details can be found in Table S3. doi:10.1371/journal.pone.0075049.g004

due to acidification alone. As the second half of the experiment is more likely to reflect the long term response of corals to ocean acidification, this negative response to OA is consistent with other studies on *Acropora* sp. [14,20,24,78], although the amount of decline differs between species. The absence of any change in calcification of *Pocillopora damicornis* is consistent with another study [14], whereas declines in calcification of 50% in *P. meandrina* were reported [37]. The lack of any change in calcification rates of *Montipora monasteriata*, and *Turbinaria reniformis* due to acidification (Fig. 2d, f, h) is in contrast to other studies which reported a 15–20% decline in *Montipora capitata* [16], and a 13% decline in *T. reniformis* (albeit at pCO₂ levels that were considerably higher than those in the present study) [78]. Although a significant interaction of pCO₂ and temperature was observed in *P. damicornis* during the first half of the experiment (Fig. 2c), this was not observed during the second half. Similarly, *M. monasteriata* calcified more at elevated compared to ambient temperature during the first half (Fig. 2e), but not during the second half of the experiment. Thus, it appears that with the exception of *A. millepora*, these species may be resistant to changes in pCO₂ and temperature within the parameter ranges investigated in this study.

In the current study, elevated temperature did not exacerbate or counteract the negative effects of OA on calcification in *A.*

millepora, and did not have an overall negative affect on calcification in the other three species. This is in contrast to other studies where elevated temperature was found to mitigate negative OA effects. For example, Anthony *et al.* [20] found that elevated temperature (28–29°C vs. 25–26°C) prevented a decline of calcification in *A. intermedia* at elevated pCO₂ (520–705 μatm). Muehllehner and Edmunds [37] showed that the negative effects of elevated pCO₂ (720 μatm) were fully mediated in *P. meandrina* when OA was combined with elevated temperature (29°C vs. 27°C). Overall, these findings add to the growing body of evidence that the response of coral calcification to OA is highly species specific, and that some coral species may maintain calcification under combined ocean acidification and warming in the future.

Although the current study was conducted using artificial seawater, it is unlikely that this influenced the observed responses of calcification to ocean acidification. The carbonate chemistry of the custom-made seawater mimicked natural conditions very well (Table 1), and calcification rates – as well as chlorophyll *a* concentrations, endosymbiont densities, energy reserves, and tissue biomass – were within the range observed in the field and/or other experimental studies using natural seawater [13,28,43,45,46, 66,78–89].

While many studies note a decline in coral calcification with increasing $p\text{CO}_2$ [75–77], there is considerable among-study variation [75,76], and some species are more resistant than others [15,90]. Such differences may be due to experimental duration, how seawater carbonate chemistry is altered (i.e., bubbling CO_2 vs. acid addition), and how calcification is measured (i.e., buoyant weight vs. total alkalinity anomaly technique). Meta-analyses have shown that experimental duration or the method of carbonate chemistry manipulation did not explain the large variability of responses observed among studies [75,76]. While this study suggests that experimental duration can influence the response of calcification to OA in some species (i.e., calcification of *A. millepora* decreased only during the second half of the second half), it is likely that biological aspects have a stronger influence on the sensitivity of coral calcification to OA than differences in methodology. Important biological aspects include energetic status and feeding [28,38], enhanced algal production [41,91], and cellular pH control [92–95].

Despite the assumption that calcification becomes energetically more costly under OA [38–40], energy reserves did not decline with increasing $p\text{CO}_2$ (Fig. 4). Lipid concentrations increased under OA conditions in both *A. millepora* and *P. damicornis*, and were fully maintained in *M. monasteriata* and *T. reniformis*. Protein, carbohydrate, and tissue biomass were overall maintained under OA conditions in all species. Further, temperature did not negatively affect energy reserves and tissue biomass except for carbohydrate concentrations in *A. millepora*, which were lower at elevated compared to ambient temperature. Importantly, energy reserves and tissue biomass were fully maintained or even increased at the highest $p\text{CO}_2$ level in *A. millepora* despite dramatic decreases in calcification rates. These findings suggest that (1) energy reserves are generally not metabolized under OA conditions or OA at elevated temperature, and (2) that either energy reserves do not play a role in sustaining calcification under OA conditions, or that the increased energetic costs of maintaining calcification under OA are relatively insignificant. This is consistent with other work showing that calcification likely does not become energetically more costly under OA conditions [81], and that the extra energy required to up-regulate pH at the site of calcification under OA conditions is <1% of that produced by photosynthesis [92].

Further, from an energetic standpoint of view, the total amount of energy reserves present in a coral species did not seem to be related to their calcification response to OA. The energetic content of lipid, protein, and carbohydrates is better assessed from an energetic point of view [96,97], as specific enthalpies of combustion differ significantly among energy reserve pools: -39.5 kJ g^{-1} for lipid, -23.9 kJ g^{-1} for protein, and -17.5 kJ g^{-1} for carbohydrate [96]. When the total amount of energy available to each species was calculated (i.e., the sum of lipid, protein, and carbohydrate expressed in kJ g^{-1} tissue biomass), *A. millepora* had the lowest amount of all species in the control treatment (6.6 vs. up to 8.1 kJ g^{-1} in *P. damicornis*), but the highest amount in the high- CO_2 treatment (9.0 vs. 6.6 kJ g^{-1} in *M. monasteriata*), and a similar amount as both *M. monasteriata* and *T. reniformis* in the high- CO_2 –high temperature treatment (7.3 kJ g^{-1} vs. 7.2 and 7.8 kJ g^{-1} , respectively). It is therefore unlikely that high levels of energy reserves *per se* help corals maintain calcification rates under OA conditions.

However, maintaining energy reserves and tissue biomass under ocean acidification does have crucial implications for other aspects of coral health and resistance to stressors such as coral bleaching. For example, maintenance of lipid concentrations may enable corals to maintain their reproductive output [98], even under

future OA and warming. This may be critical considering that many other processes involved in coral reproduction such as fertilization, settlement success, and metamorphosis are compromised under OA [99,100]. Furthermore, maintenance of energy reserves has been shown to be associated with higher resistance to coral bleaching and to promote recovery from bleaching [46,49], which could prove critical as bleaching events will increase in frequency over the coming decades [101].

Heterotrophy is known to promote energy storage, tissue synthesis, and skeletal growth in healthy and bleached corals [47,48,102] as well as corals subjected to OA [28,42]. Therefore, biweekly feeding in this study (intended to mimic zooplankton contribution to the coral diet on the reef) may have helped corals to sustain energy reserves and tissue biomass under these conditions. It has further been suggested that coral tissue reacts to availability of such resources faster than skeletal growth [103,104], which could explain why tissue biomass – but not necessarily calcification – was maintained or even increased in all four species irrespective of $p\text{CO}_2$ or temperature conditions. As feeding rates and heterotrophic plasticity are highly species-specific [48,54,105], it is likely that heterotrophic carbon intake differed significantly among the species studied here, potentially contributing to their differential responses to OA.

Enhanced algal productivity due to CO_2 -fertilization [41,91] may help corals to maintain calcification under OA conditions. Although chlorophyll *a* concentrations and endosymbiont density were unaffected at the highest $p\text{CO}_2$ level (except for chlorophyll in *T. reniformis*), CO_2 -fertilization may nevertheless have played a role in helping corals to maintain energy reserves and/or calcification. Increased availability of $\text{CO}_{2(\text{aq})}$ under OA conditions may enhance algal productivity, especially in *Symbiodinium* phylotypes with less efficient carbon-concentrating mechanisms, which rely to a greater extent on the passive, diffusive uptake of $\text{CO}_{2(\text{aq})}$ [41]. Thus, a potentially increased translocation of autotrophic carbon to the animal host may have contributed to the maintenance of energy reserves and tissue biomass observed here.

Interestingly, both chlorophyll *a* concentrations and endosymbiont density were often lowest at $607 \mu\text{atm}$, showing a non-linear relationship with increasing $p\text{CO}_2$. Nevertheless, the lack of any significant difference in chlorophyll *a* and/or symbiont density at $741 \mu\text{atm}$ versus ambient $p\text{CO}_2$ concentrations (except for chlorophyll in *T. reniformis*) is consistent with other studies [23,28,29,106]. The reason for the observed minima at $\sim 600 \mu\text{atm}$ is unknown. Similar non-linear responses were not observed for calcification rates, tissue biomass, and most energy reserve pools, suggesting that this did not translate into a decreased performance of the animal host. Edmunds [81] also observed a non-linear $p\text{CO}_2$ threshold between 756 and $861 \mu\text{atm}$ affecting photochemistry and respiration in massive *Porites* corals, thus highlighting the importance of studying multiple $p\text{CO}_2$ levels in OA experiments in order to assess non-linear physiological responses and to better forecast physiological responses over the coming century as the oceans continue to warm and acidify.

In addition to energetic status and enhanced algal productivity due to CO_2 fertilization, other factors such as the amount of control over the carbonate chemistry at the site of calcification may explain the observed differences in susceptibility of calcification to OA. Corals have the ability to significantly up-regulate the pH at the site of calcification compared to ambient seawater, even under OA conditions [92–95,107,108]. Yet, the degree to which corals are able to control the pH at the site of calcification likely varies among species [92,95,109]. *Acropora* spp. may have the lowest capacity to up-regulate pH at the site of calcification based

on boron isotopic measurements [92,110]. Further, crystallization under OA was most compromised in *Acropora verweyi* and least compromised in *T. reniformis* [78]. Thus, we hypothesize that *A. millepora* has a weaker proton pump than the other coral species studied here, making its calcification rate more sensitive to future OA. Although pH up-regulation has not been studied in *P. damicomis*, *M. monasteriata*, or *T. reniformis*, it can be hypothesized that they have stronger control over the pH at the site of calcification and were therefore able to maintain calcification under the $p\text{CO}_2$ levels studied here.

As physiological responses of both the animal host and algal endosymbiont to combined OA and warming were strongly species-specific, a wide range of susceptibility patterns can be expected resulting in ecological “winners and losers” [90,111]. Branching *Acropora* corals, which are important reef builders, are likely to be “losers” on future coral reefs because they are highly susceptible to both bleaching [111] and OA [20,24,90,99]. This can be expected to have severe impacts on reef diversity, structural complexity, and overall reef functioning. Nevertheless, some corals could be more resistant to combined ocean acidification and warming expected by the end of this century than previously thought, as three of the four species fully maintained calcification under elevated $p\text{CO}_2$ and temperature without compromising overall energy reserves or biomass. Further, the immediate effects of rising seawater temperature and ocean acidification may be tolerable for some species.

References

- Hoegh-Guldberg O, Mumby PJ, Hooten AJ, Steneck R, Greenfield P, et al. (2007) Coral reefs under rapid climate change and ocean acidification. *Science* 318: 1737–1742.
- Hoegh-Guldberg O, Bruno JF (2010) The impact of climate change on the world's marine ecosystems. *Science* 328: 1523–1528.
- IPCC (2007) Climate Change 2007: The physical science basis. Summary for policymakers. Contribution of working group I to the fourth assessment report of the Intergovernmental Panel on Climate Change. <http://www.ipcc.ch> website. Accessed 2013 August 19.
- Sabine CL, Feely RF, Gruber N, Key RM, Lee K, et al. (2004) The oceanic sink for anthropogenic CO_2 . *Science* 305: 367–371.
- Caldeira K, Wickett ME (2003) Anthropogenic carbon and ocean pH. *Nature* 425: 365.
- Fitt WK, Brown BE, Warner ME, Dunne RP (2001) Coral bleaching: interpretation of thermal tolerance limits and thermal thresholds in tropical corals. *Coral Reefs* 20: 51–65.
- Wild C, Hoegh-Guldberg O, Naumann MS, Colombo-Pallotta MF, Atewberhan M, et al. (2011) Climate change impedes scleractinian corals as primary reef ecosystem engineers. *Marine and Freshwater Research* 62: 205–215.
- Kleypas J, Buddemeier RW, Archer D, Gattuso J-P, Langdon C, et al. (1999) Geochemical consequences of increased atmospheric carbon dioxide on coral reefs. *Science* 284: 118–120.
- Feely RA, Sabine CL, Lee K, Berelson W, Kleypas J, et al. (2004) Impact of anthropogenic CO_2 on the CaCO_3 system in the oceans. *Science* 305: 362–366.
- Orr JC, Fabry VJ, Aumont O, Bopp L, Doney SC, et al. (2005) Anthropogenic ocean acidification over the twenty-first century and its impact on calcifying organisms. *Nature* 437: 681–686.
- Feely RA, Doney SC, Cooley SR (2009) Ocean acidification: present condition and future changes in a high- CO_2 world. *Oceanography* 22: 36–47.
- Langdon C, Takahashi T, Sweeney C, Chipman D, Goddard J, et al. (2000) Effect of calcium carbonate saturation state on the calcification rate of an experimental coral reef. *Global Biogeochemical Cycles* 14: 639–654.
- Krief S, Hendy EJ, Fine M, Yam R, Meibom A, et al. (2010) Physiological and isotopic responses of scleractinian corals to ocean acidification. *Geochimica et Cosmochimica Acta* 74: 4988–5001.
- Comeau S, Edmunds PJ, Spindel NB, Carpenter RC (2013) The responses of eight coral reef calcifiers to increasing partial pressure of CO_2 do not exhibit a tipping point. *Limnology and Oceanography* 58: 388–398.
- Edmunds PJ, Brown D, Moriarty V (2012) Interactive effects of ocean acidification and temperature on two scleractinian corals from Moorea, French Polynesia. *Global Change Biology* 18: 2173–2183.
- Jokiel PL, Rodgers KS, Kuffner IB, Andersson AJ, Cox EF, et al. (2008) Ocean acidification and calcifying reef organisms: a mesocosm investigation. *Coral Reefs* 27: 473–483.

Supporting Information

Table S1 Results of 8 two-way ANOVAs for average calcification rate during the first and second half of the experiment. (DOCX)

Table S2 Results of 8 two-way ANOVAs for average chlorophyll a concentrations and symbiont density. (DOCX)

Table S3 Results of 16 two-way ANOVAs for average soluble lipid, animal soluble protein, animal soluble carbohydrate concentrations, and tissue biomass. (DOCX)

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Author Contributions

Conceived and designed the experiments: AGG MEW WJC. Performed the experiments: VS AGG MEW WJC TFM KDH DTP XH QL HX YW YM JHB. Analyzed the data: VS. Contributed reagents/materials/analysis tools: VS AGG MEW WJC TFM KDH DTP XH QL HX YW YM JHB. Wrote the paper: VS AGG MEW WJC TFM KDH DTP XH QL HX YW YM JHB.

- Leclercq N, Gattuso J-P, Jaubert J (2000) CO_2 partial pressure controls the calcification rate of a coral community. *Global Change Biology* 6: 329–334.
- Renegar DA, Riegl B (2005) Effect of nutrient enrichment and elevated CO_2 partial pressure on growth rate of Atlantic scleractinian coral *Acropora cervicornis*. *Marine Ecology Progress Series* 293: 69–76.
- Langdon C, Atkinson MJ (2005) Effect of elevated $p\text{CO}_2$ on photosynthesis and calcification on corals and interactions with seasonal change in temperature/irradiance and nutrient enrichment. *Journal of Geophysical Research* 110: C09S07, doi:10.1029/2004JC002576
- Anthony KRN, Kline DI, Diaz-Pulido G, Dove S, Hoegh-Guldberg O (2008) Ocean acidification causes bleaching and productivity loss in coral reef builders. *Proceedings of the National Academy of Sciences of the USA* 105: 17442–17446.
- Holcomb M, Cohen AL, McCorkle DC (2012) An investigation of the calcification response of the scleractinian corals *Astrangia poculata* to elevated $p\text{CO}_2$ and the effects of nutrients, zooxanthellae and gender. *Biogeosciences* 9: 29–39.
- Marubini F, Atkinson MJ (1999) Effects of lowered pH and elevated nitrate on coral calcification. *Marine Ecology Progress Series* 188: 117–121.
- Marubini F, Ferrier-Pages C, Furla P, Allemand D (2008) Coral calcification responds to seawater acidification: a working hypothesis towards a physiological mechanism. *Coral Reefs* 27: 491–499.
- Schneider K, Erez J (2006) The effect of carbonate chemistry on calcification and photosynthesis in the hermatypic coral *Acropora eurystoma*. *Limnology and Oceanography* 51: 1284–1293.
- Jury CP, Whitehead RF, Szmant AM (2010) Effects of variation in carbonate chemistry on the calcification rates of *Madraca auretenra* (*Madraca mirabilis* sensu Wells, 1973): bicarbonate concentrations best predict calcification rates. *Global Change Biology* 16: 1632–1644.
- Comeau S, Carpenter R, Edmunds P (2013) Effects of feeding and light intensity on the response of the coral *Porites rus* to ocean acidification. *Marine Biology* 160: 1127–1134.
- Rodolfo-Metalpa R, Houlbreque F, Tambutte E, Boisson F, Baggini C, et al. (2011) Coral and mollusc resistance to ocean acidification adversely affected by warming. *Nature Climate Change* 1: 308–312.
- Edmunds PJ (2011) Zooplanktivory ameliorates the effects of ocean acidification on the reef coral *Porites* spp. *Limnology and Oceanography* 56: 2402–2410.
- Houlbreque F, Rodolfo-Metalpa R, Jeffrey R, Oberhaensli F, Teyssie J-L, et al. (2012) Effects of increased $p\text{CO}_2$ on zinc uptake and calcification in the tropical coral *Stylophora pistillata*. *Coral Reefs* 31: 101–109.
- Reynaud S, Leclercq N, Romaine-Lioud S, Ferrier-Pages C, Jaubert J, et al. (2003) Interacting effects of CO_2 partial pressure and temperature on photosynthesis and calcification in a scleractinian coral. *Global Change Biology* 9: 1660–1668.

31. Rodolfo-Metalpa R, Martin S, Ferrier-Pages C, Gattuso J-P (2010) Response of the temperate coral *Cladocora caespitosa* to mid- and long-term exposure to pCO₂ and temperature levels projected for the year 2100 AD. *Biogeosciences* 7: 289–300.
32. Clausen C (1971) Effects of temperature on the rate of ⁴⁵Ca uptake by *Pocillopora damicornis*. In: Lenhoff HM, Muscatine L, Davis LV, editors. *Experimental Coelenterate Biology*. Honolulu: University of Hawaii Press. pp. 246–260.
33. Cantin NE, Cohen AL, Karnauskas KB, Tarrant AM, McCorkle DC (2010) Ocean warming slows coral growth in the Central Red Sea. *Science* 329: 322–325.
34. Clausen C, Roth AA (1975) Effect of temperature and temperature adaptation on calcification rates in the hermatypic coral *Pocillopora damicornis*. *Marine Biology* 33: 93–100.
35. Jokiel PL, Coles SL (1977a) Effects of temperature on the mortality and growth of Hawaiian reef corals. *Marine Biology* 43: 201–208.
36. Marshall AT, Clode P (2004) Calcification rate and the effect of temperature in a zooxanthellate and an azooxanthellate scleractinian reef coral. *Coral Reefs* 23: 218–224.
37. Muehllehner N, Edmunds PJ (2008) Effects of ocean acidification and increased temperature on skeletal growth of two scleractinian corals, *Pocillopora meandrina* and *Porites* sp. Ft. Lauderdale, Florida, 7–11 July 2008: Proceedings of the 11th International Coral Reef Symposium. 57–61p.
38. Cohen AL, Holcomb M (2009) Why corals care about ocean acidification - uncovering the mechanism. *Oceanography* 22: 118–127.
39. Erez J, Reynaud S, Silverman J, Schneider K, Allemand D (2011) Coral calcification under ocean acidification and global change. In: Stambler N, Dubinsky Z, editors. *Coral Reefs: An Ecosystem in Transition*: Springer. 552p.
40. Pandolfi JM, Connolly SR, Marshall DJ, Cohen AL (2011) Projecting coral reef futures under global warming and ocean acidification. *Science* 333: 418–422.
41. Brading P, Warner ME, Davey P, Smith DJ, Achterberg EP, et al. (2011) Differential effects of ocean acidification on growth and photosynthesis among phylogenetic types of *Symbiodinium* (Dinophyceae). *Limnology and Oceanography* 56: 927–938.
42. Drenkard EJ, Cohen AL, McCorkle DC, Putron SJ, Starczak VR, et al. (2013) Calcification by juvenile corals under heterotrophy and elevated CO₂. *Coral Reefs*: doi:10.1007/s00338-0013-01021-00335
43. Fitt WK, McFarland FK, Warner ME, Chilcoat GC (2000) Seasonal patterns of tissue biomass and densities of symbiotic dinoflagellates in reef corals and relation to coral bleaching. *Limnology and Oceanography* 45: 677–685.
44. Stimson J (1997) The annual cycle of density of zooxanthellae in the tissue of field and laboratory-held *Pocillopora damicornis* (Linnaeus). *Journal of Experimental Marine Biology and Ecology* 214: 35–48.
45. Thornhill DJ, Rotjan RD, Todd BD, Chilcoat GC, Iglesias-Prieto R, et al. (2011) A connection between colony biomass and death in Caribbean reef-building corals. *PLoS ONE* 6: e29535. doi:10.1371/journal.pone.0029535
46. Rodrigues LJ, Grottoli AG (2007) Energy reserves and metabolism as indicators of coral recovery from bleaching. *Limnology and Oceanography* 52: 1874–1882.
47. Levas SJ, Grottoli AG, Hughes AD, Osburn CL, Matsui Y (2013) Physiological and biogeochemical traits of bleaching and recovery in the mounding species of coral *Porites lobata*: Implications for resilience in mounding corals. *PLoS ONE* 8: e63267. doi:10.1371/journal.pone.0063267
48. Grottoli AG, Rodrigues LJ, Palardy JE (2006) Heterotrophic plasticity and resilience in bleached corals. *Nature* 440: 1186–1189.
49. Anthony KRN, Hoogenboom MO, Maynard JF, Grottoli AG, Middlebrook R (2009) Energetics approach to predicting mortality risk from environmental stress: a case study of coral bleaching. *Functional Ecology* 23: 539–550.
50. Fine M, Tchernov D (2007) Scleractinian coral species survive and recover from decalcification. *Science* 315: 1811.
51. Muscatine L, McCloskey LR, Marian RE (1981) Estimating the daily contribution of carbon from zooxanthellae to coral animal respiration. *Limnology and Oceanography* 26: 601–611.
52. Hoegh-Guldberg O (1999) Climate change, coral bleaching and the future of the world's coral reefs. *Marine and Freshwater Research* 50: 839–866.
53. Hurlbert SH (1984) Pseudoreplication and the Design of Ecological Field Experiments. *Ecological Monographs* 54: 187–211.
54. Palardy JE, Rodrigues LJ, Grottoli AG (2008) The importance of zooplankton to the daily metabolic carbon requirements of healthy and bleached corals at two depths. *Journal of Experimental Marine Biology and Ecology* 367: 180–188.
55. Grottoli AG (2002) Effect of light and brine shrimp levels on skeletal $\delta^{13}\text{C}$ values in the Hawaiian coral *Porites compressa*: A tank experiment. *Geochimica et Cosmochimica Acta* 66: 1955–1967.
56. Wang ZA, Cai W-J (2004) Carbon dioxide degassing and inorganic carbon export from a marsh-dominated estuary (the Duplin River): a Marsh CO₂ pump. *Limnology and Oceanography* 49: 341–354.
57. Cai W-J, Hu X, Huang W-J, Jian L-Q, Peng T-H, et al. (2010) Alkalinity distribution in the western North Atlantic Ocean margins. *Journal of Geophysical Research* 115: C08014.
58. Lewis E, Wallace D (1998) Program Developed for CO₂ System Calculations. ORNL/CDIAC-105. Oak Ridge, Tennessee: Carbon Dioxide Information Analysis Center, Oak Ridge National Laboratory, U.S. Department of Energy.
59. Weiss RF, Bullister JL, Gammon RH, Warner MJ (1985) Atmospheric chlorofluoromethanes in the deep equatorial Atlantic. *Nature* 314: 608–610.
60. Millero FJ, Graham TB, Huang F, Bustos-Serrano H, Pierrot D (2006) Dissociation constants of carbonic acid in seawater as a function of salinity and temperature. *Marine Chemistry* 100: 80–94.
61. Jokiel PL, Maragos JE, Franzisket L (1978) Coral growth: buoyant weight technique. In: Stoddart DR, Johannes RE, editors. *Coral Reefs: Research Methods*. Paris: UNESCO. 529–541.
62. Johannes RE, Wiebe WJ (1970) A method for determination of coral tissue biomass and composition. *Limnology and Oceanography* 21: 540–547.
63. Porra RJ, Tompson WA, Kriedemann PE (1989) Determination of accurate extinction coefficients and simultaneous equations for assaying chlorophylls a and b extracted with four different solvents: verification of the concentration of chlorophyll standards by atomic absorption spectroscopy. *Biochimica et Biophysica Acta* 975: 384–394.
64. Warren CR (2008) Rapid measurement of chlorophylls with a microplate reader. *Journal of Plant Nutrition* 31: 1321–1332.
65. Oku H, Yamashiro H, Onaga K, Iwasaki H, Takara K (2002) Lipid distribution in branching coral *Montipora digitata*. *Fisheries Science* 68: 517–522.
66. Grottoli AG, Rodrigues LJ, Juarez C (2004) Lipids and stable carbon isotopes in two species of Hawaiian corals, *Porites compressa* and *Montipora verrucosa*, following a bleaching event. *Marine Biology* 145: 621–631.
67. Smith PK, Krohn RI, Hermanson GT, Mallia AK, Gartner FH, et al. (1985) Measurement of protein using bicinchoninic acid. *Analytical Biochemistry* 150: 76–85.
68. Dubois M, Giles KA, Hamilton JK, Pebers PA, Smith F (1956) Colorimetric method for determination of sugar and related substances. *Analytical Chemistry* 28: 350–356.
69. Marsh JA (1970) Primary productivity of reef-building calcareous red algae. *Ecology* 51: 255–263.
70. Stimson J, Kinzie RA (1991) The temporal pattern and rate of release of zooxanthellae from the reef coral *Pocillopora damicornis* (Linnaeus) under nitrogen-enrichment and control conditions. *Journal of Experimental Marine Biology and Ecology* 153: 63–74.
71. Veal CJ, Carmi M, Fine M, Hoegh-Guldberg O (2010) Increasing the accuracy of surface area estimation using single wax dipping of coral fragments. *Coral Reefs* 29: 893–897.
72. Winer BJ (1971) *Statistical Principles in Experimental Design*. New York: McGraw-Hill.
73. Quinn GP, Keough MJ (2002) *Experimental Design and Data Analysis for Biologists*. New York: Cambridge University Press, 557p.
74. Moran MD (2003) Arguments for rejecting the sequential Bonferroni in ecological studies. *Oikos* 100: 403–405.
75. Chan NCS, Connolly SR (2012) Sensitivity of coral calcification to ocean acidification: a meta-analysis. *Global Change Biology*: doi:10.1111/gcb.12011
76. Kroeker KJ, Kordas RL, Crim RN, Singh GG (2010) Meta-analysis reveals negative yet variable effects of ocean acidification on marine organisms. *Ecology Letters* 13: 1419–1434.
77. Harvey BP, Gwynn-Jones D, Moore PJ (2013) Meta-analysis reveals complex marine biological responses to the interactive effects of ocean acidification and warming. *Ecology and Evolution*: doi:10.1002/ece1003.1516
78. Marubini F, Ferrier-Pages C, Cuif J-P (2003) Suppression of skeletal growth in scleractinian corals by decreasing ambient carbonate-ion concentration: a cross-family comparison. *Proceedings of the Royal Society B* 270: 179–184.
79. D'Croz L, Mate JL (2004) Experimental responses to elevated water temperature in genotypes of the reef coral *Pocillopora damicornis* from upwelling and non-upwelling environments in Panama. *Coral Reefs* 23: 473–483.
80. Reynaud-Vaganay S, Juillet-Leclerc A, Jaubert J, Gattuso J-P (2001) Effect of light on skeletal $\delta^{13}\text{C}$ and $\delta^{18}\text{O}$, and interaction with photosynthesis, respiration and calcification in two zooxanthellate scleractinian corals. *Palaeogeography, Palaeoclimatology, Palaeoecology* 175: 393–403.
81. Edmunds PJ (2012) Effect of pCO₂ on the growth, respiration, and photophysiology of massive *Porites* spp. in Moorea, French Polynesia. *Marine Biology* 159: 2149–2160.
82. Dove S, Ortiz J-C, Enriquez S, Fine M, Fisher P, et al. (2006) Response of holosymbiont pigments from the scleractinian coral *Montipora monasteriata* to short-term heat stress. *Limnology and Oceanography* 51: 1149–1158.
83. Ferrier-Pages C, Rottier C, Beraud E, Levy O (2010) Experimental assessment of the feeding effort of three scleractinian coral species during a thermal stress: Effect on the rates of photosynthesis. *Journal of Experimental Marine Biology and Ecology* 390: 118–124.
84. Levas SJ (2012) *Biogeochemistry and Physiology of Bleached and Recovering Hawaiian and Caribbean Corals*. Columbus, OH: PhD Thesis, The Ohio State University. 238p.
85. Muller-Parker G, McCloskey LR, Hoegh-Guldberg O, McAuley PJ (1994) Effect of ammonium enrichment on animal and algal biomass of the coral *Pocillopora damicornis*. *Pacific Science* 48: 273–283.
86. Stambler N, Popper N, Dubinsky Z, Stimson J (1991) Effects of nutrient enrichment and water motion on the coral *Pocillopora damicornis*. *Pacific Science* 45: 299–307.
87. Takabayashi M, Hoegh-Guldberg O (1995) Ecological and physiological differences between two colour morphs of the coral *Pocillopora damicornis*. *Marine Biology* 123: 705–714.

88. Tolosa I, Treignier C, Grover R, Ferrier-Pages C (2011) Impact of feeding and short-term temperature stress on the content and isotopic signature of fatty acids, sterols, and alcohols in the scleractinian coral *Turbinaria reniformis*. *Coral Reefs* 30: 763–774.
89. Treignier C, Grover R, Ferrier-Pages C, Tolosa I (2008) Effect of light and feeding on the fatty acid and sterol composition of zooxanthellae and host tissue isolated from the scleractinian coral *Turbinaria reniformis*. *Limnology and Oceanography* 53: 2702–2710.
90. Fabricius KE, Langdon C, Uthicke S, Humphrey C, Noonan S, et al. (2011) Losers and winners on coral reefs acclimatized to elevated carbon dioxide concentrations. *Nature Climate Change* 1: 165–169.
91. Herfort L, Thake B, Taubner I (2008) Bicarbonate stimulation of calcification and photosynthesis in two hermatypic corals. *Journal of Phycology* 44: 91–98.
92. McCulloch MT, Falter J, Trotter J, Montagna P (2012) Coral resilience to ocean acidification and global warming through pH up-regulation. *Nature Climate Change*: doi:10.1038/nclimate1473
93. Venn A, Tambutte E, Holcomb M, Allemand D, Tambutte S (2011) Live tissue imaging shows reef corals elevate pH under their calcifying tissue relative to seawater. *PLoS ONE* 6: e20013.
94. Venn AA, Tambutte E, Holcomb M, Laurent J, Allemand D, et al. (2013) Impact of seawater acidification on pH at the tissue-skeleton interface and calcification in reef corals. *Proceedings of the National Academy of Sciences* 110: 1634–1639.
95. Ries JB (2011) A physicochemical framework for interpreting the biological calcification response to CO₂-induced ocean acidification. *Geochimica and Cosmochimica Acta* 75: 4053–4064.
96. Gnaiger E, Bitterlich G (1984) Proximate biochemical composition and caloric content calculated from elemental CHN analysis: a stoichiometric concept. *Oecologia* 62: 289–298.
97. Lesser MP (2013) Using energetic budgets to assess the effects of environmental stress on corals: are we measuring the right things? *Coral Reefs* 32: 25–33.
98. Ward S (1995) Two patterns of energy allocation for growth, reproduction and lipid storage in the scleractinian coral *Pocillopora damicornis*. *Coral Reefs* 14: 87–90.
99. Albright R, Mason B, Miller M, Langdon C (2010) Ocean acidification compromises recruitment success of the threatened Caribbean coral *Acropora palmata*. *Proceedings of the National Academy of Sciences of the USA* 107: 20400–20404.
100. Nakamura M, Ohki S, Suzuki A, Sakai K (2011) Coral larvae under ocean acidification: survival, metabolism, and metamorphosis. *Public Library of Science One* 6: e14521. doi:10.1371/journal.pone.0014521
101. Donner SD (2009) Coping with commitment: Projected thermal stress on coral reefs under different future scenarios. *PLoS ONE* 4: e5712.
102. Houlbreque F, Ferrier-Pages C (2009) Heterotrophy in tropical scleractinian corals. *Biological Reviews* 84: 1–17.
103. Houlbreque F, Tambutte E, Allemand D, Ferrier-Pages C (2004) Interactions between zooplankton feeding, photosynthesis and skeletal growth in the Scleractinian coral *Stylophora pistillata*. *Journal of Experimental Biology* 207: 1461–1469.
104. Anthony KRN, Connolly SR, Willis BL (2002) Comparative analysis of energy allocation to tissue and skeletal growth in corals. *Limnology and Oceanography* 47: 1417–1429.
105. Palardy JE, Grotto AG, Matthews KA (2005) Effects of upwelling, depth, morphology and polyp size on feeding of three species of Panamanian corals. *Marine Ecology Progress Series* 300: 79–89.
106. Godinot C, Houlbreque F, Grover R, Ferrier-Pages C (2011) Coral uptake of inorganic phosphorus and nitrogen negatively affected by simultaneous changes in temperature and pH. *PLoS ONE* 6: e25024.
107. Al-Horani FA, Al-Moghrabi SM, de Beer D (2003) The mechanism of calcification and its relation to photosynthesis and respiration in the scleractinian coral *Galaxea fascicularis*. *Marine Biology* 142: 419–426.
108. Cohen AL, McCorkle DC, de Putron S, Gaetani GA, Rose KA (2009) Morphological and compositional changes in the skeletons of new coral recruits reared in acidified seawater: Insights into the biomineralization response to ocean acidification. *Geochemistry, Geophysics, Geosystems* 10: Q07005, doi:07010.01029/02009GC002411
109. Ries JB, Cohen AL, McCorkle DC (2009) Marine calcifiers exhibit mixed responses to CO₂-induced ocean acidification. *Geology* 37: 1131–1134.
110. Trotter J, Montagna P, McCulloch MT, Silenzi S, Reynaud S, et al. (2011) Quantifying the pH 'vital effect' in the temperate zooxanthellate coral *Cladocora caespitosa*: Validation of the boron seawater pH proxy. *Earth and Planetary Science Letters* 303: 163–173.
111. Loya Y, Sakai K, Yamazato K, Nakan Y, Sambali H, et al. (2001) Coral bleaching: the winners and the losers. *Ecology Letters* 4: 122–131.